

The ARM SCM Intercomparison Study—Overview and Preliminary Results for Case 1

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Introduction

The Single-Column Model (SCM) Working Group in the Atmospheric Radiation Measurement (ARM) Program has begun a series of SCM Intercomparison case studies to evaluate the adequacy of the forcing data sets and the progress of SCM formulations. There are nine modeling groups participating, which include eight SCMs: Ghan [Pacific Northwest National Laboratory (PNNL)], Randall/Cripe [Colorado State University (CSU)], Somerville/Iacobellis [Scripps/University of California, San Diego (UCSD)], Klein [National Oceanic and Atmospheric Administration (NOAA)-Geophysical Fluid Dynamics Laboratory (GFDL)], Lohmann (Dalhousie), Stenchikov/Robock (Maryland, Rutgers), Zhang/Xie [State University of New York (SUNY) Stony Brook], Sud/Wlaker [National Aeronautics and Space Administration (NASA)-Goddard Space Flight Center (GSFC)], and one two-dimensional (2-D) cloud-resolving model: Xu (CSU). The first case study addresses the prescription of advective forcing, the methods used to derive SCM forcing terms, and the methods used to estimate surface flux forcing. Details of the SCM Intercomparison procedures are given elsewhere (Cederwall and Krueger 1998).

Approach

A critical issue in single-column modeling is the prescription of advective forcing, required for the SCMs because the information usually provided by neighboring cells in a GCM is not available in the single-column formulation. The SCM large-scale (L.S.) advective forcing comprises horizontal and vertical advection, given here for water vapor mixing ratio (q) as an example:

$$\left(\frac{\partial \bar{q}}{\partial t}\right)_{L.S.} \equiv -\bar{\mathbf{v}} \cdot \nabla \bar{q} - \bar{\omega} \frac{\partial \bar{q}}{\partial p}$$

Three methods for specifying the advecting forcing were tested, defined by the following expressions (given for q):

1. observed total advective tendency,

$$\left(\frac{\partial \bar{q}}{\partial t}\right)_{L.S.} = \left(\frac{\partial \bar{q}}{\partial t}\right)_{L.S.T.}$$

2. observed horizontal advective tendency plus vertical advective tendency estimated using observed large-scale vertical motion and the model-predicted vertical gradient,

$$\left(\frac{\partial \bar{q}}{\partial t}\right)_{L.S.} = \left(\frac{\partial \bar{q}}{\partial t}\right)_{L.S.H.} - \bar{\omega} \frac{\partial \bar{q}_m}{\partial p}$$

3. horizontal advective tendency estimated using a relaxation toward upstream values, plus vertical advective tendency estimated as in Eq. (2) above,

$$\left(\frac{\partial \bar{q}}{\partial t}\right)_{L.S.} = -\frac{\bar{q}_m - \bar{q}_u}{\tau_{adv}} - \bar{\omega} \frac{\partial \bar{q}_m}{\partial p}$$

SCM advective tendency terms were obtained in two ways: 1) Barnes objective analysis using ARM sounding and NOAA wind profiler data, and 2) variational analysis (provided by Zhang) that uses ARM sounding and NOAA wind profiler data and then additional data to adjust the advective tendencies in order to match the observed

column-integrated tendencies of mass, moisture, static energy, and momentum (Zhang and Lin 1997).

Surface forcing was prescribed from two methods of heat and moisture flux estimates:

1. area-averaged SiB2 model output (from a 6.25-km grid) that uses ARM observations as input (Doran et al. 1998)
2. area-averaged observations from Energy Balance Bowen Ratio (EBBR) stations.

The EBBR stations are located only in non-cropland areas, and therefore sample just a part of the SCM bottom surface. The SiB2 model approach incorporates all surface types in the SCM domain. The difference in estimates of heat and moisture flux is most pronounced in clear-sky daytime periods, when the harvested wheat fields (the dominant cropland) are *hotter and drier* than the non-cropland (Shaw et al. 1998). This is illustrated in Figure 1, where Julian days 208-211 have clear skies.

The SCM Intercomparison involved an abbreviated 3 x 2 matrix of runs to evaluate the three methods of prescribing the advective forcing, the two methods of deriving the SCM forcing terms, and the two methods of estimating the surface forcing. This matrix and associated simulation notation is given in Table 1.

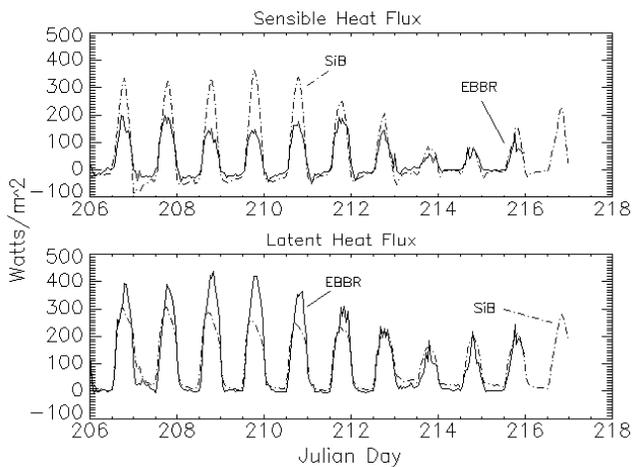


Figure 1. Estimates of (a) sensible and (b) latent heat flux from average EBBR data (solid lines) and SiB2 model output (dashed-dot lines). Clear skies occurred on Julian days 208-211.

	Barnes Analysis (LLNL)	Variational Analysis (SUNY-SB)
Observed total advective tendency	A ^(a)	D ^(a)
Observed horizontal advective tendency	B	E ^(a)
Relaxation toward upstream values	C	F ^(a)
(a) Simulations made with both surface forcing methods.		

Meteorological Conditions

The first case study is based on data from the Summer 1995 SCM Intensive Observation Period (IOP), July 18 to August 3, 1995. The first half of the period was characterized by variable cloudiness and precipitation every other day associated with a stationary, large-scale upper-level trough over North America. In the second half of the period, upper-level ridging led to clear days and hot, dry conditions. An upper-level trough replaced the ridge, with increasing cloudiness, thunderstorms, and occasional intense precipitation toward the end of the study period. Hence, a wide range of summertime weather conditions occurred for testing SCMs, including hot, clear days, variable cloudiness and local convection, and synoptic forcing with increased cloudiness, precipitation, and occasional severe weather. Inspection of an animation of the satellite images for the IOP reveals the importance of cloud advection through the study area. Cloud processes are not governed simply by local convection.

Results

Several comparisons between simulated and observed values were made for such quantities as temperature and moisture profiles, surface and top-of-atmosphere (TOA) radiative fluxes, column-integrated cloud-liquid water, precipitable water, and rainfall rate. A few of these comparisons are illustrated next.

A comparison is made to evaluate the benefit of deriving SCM forcing terms with variational analysis over those derived from traditional Barnes objective analysis. Profiles of model bias (model - observed values) for water vapor mixing ratio are shown for seven of the participating models using Barnes analysis (simulation B) and variational

analysis (simulation E). For simulation B, there is clearly a dry bias in the lower troposphere (see Figure 2a). This bias is removed in simulation E (see Figure 2b). The use of integrated water vapor from the ARM microwave radiometers in the variational analysis helps adjust the analyzed values to the state of the atmosphere and provides more representative advective forcing for the SCMs.

The ability of the model to simulate the state of the atmosphere and the radiative interactions is indicated by the outgoing longwave radiation. A comparison is made for variational-analysis derived SCM terms between the ‘total advective tendency’ prescription of advective forcing and the ‘relaxation toward upwind values’ prescription. An illustration of this comparison is given in Figure 3 for three of the SCMs. Generally, the ‘relaxation’ mode gives slightly better simulations than the ‘total’ mode (see Figure 3c vs. Figure 3b). In clear-sky conditions (Julian days 208-211), both advection prescriptions produce simulations with reduced model bias compared to those in cloudy-sky conditions. This is not surprising, and confirms the challenges that remain for improving parameterizations of clouds and their radiative interactions.

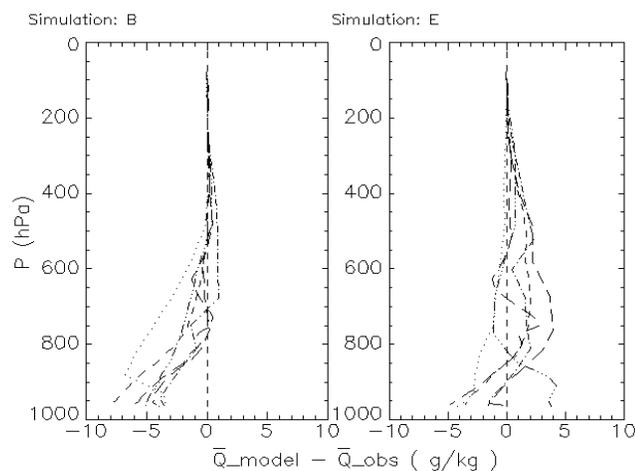


Figure 2. Vertical profiles of model bias of simulated water vapor mixing ratio for seven models using (a) Barnes objective analysis in simulation B, and (b) variational analysis in simulation E.

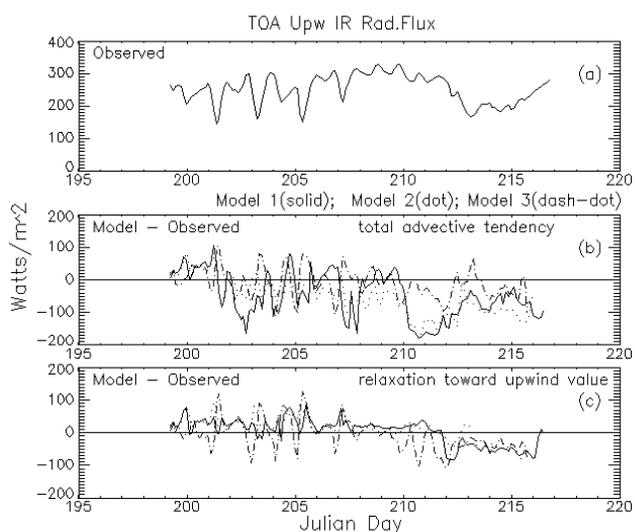


Figure 3. Time series of outgoing longwave radiation (W/m^2) from (a) satellite-based observations, (b) model bias for simulation D—‘total advective tendency,’ and (c) model bias for simulation F—‘relaxation toward upwind value.’ Model results are for three SCMs.

Preliminary Conclusions

Based on the preliminary analyses for the Case 1 simulations, we have drawn the following, tentative conclusions:

1. Simulations are improved with forcing terms derived by the *variational analysis*.
2. Simulations match observations better in *clear-sky* conditions than in *cloudy* conditions.
3. The *relaxation* toward upstream values is the preferred prescription for advective tendency when evaluating *process parameterizations*, especially those for clouds.
4. In general, the *2-D CRM* performs better than the SCMs.

The results were mixed for simulations using the two different estimates of surface forcing. This is under further study by a subgroup of participants.

Future Efforts

The SCM Working Group is conducting more SCM Intercomparison case studies. Another summertime case is planned, in collaboration with the Global Energy and Water Experiment (GEWEX) Cloud System Study Working Group 4, studying deep convection. Data from the Summer 1997 SCM IOP will be used. A fall case, using the Fall 1997 SCM IOP, is also planned to take advantage of the five other ARM IOPs that were conducted at that time.

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