

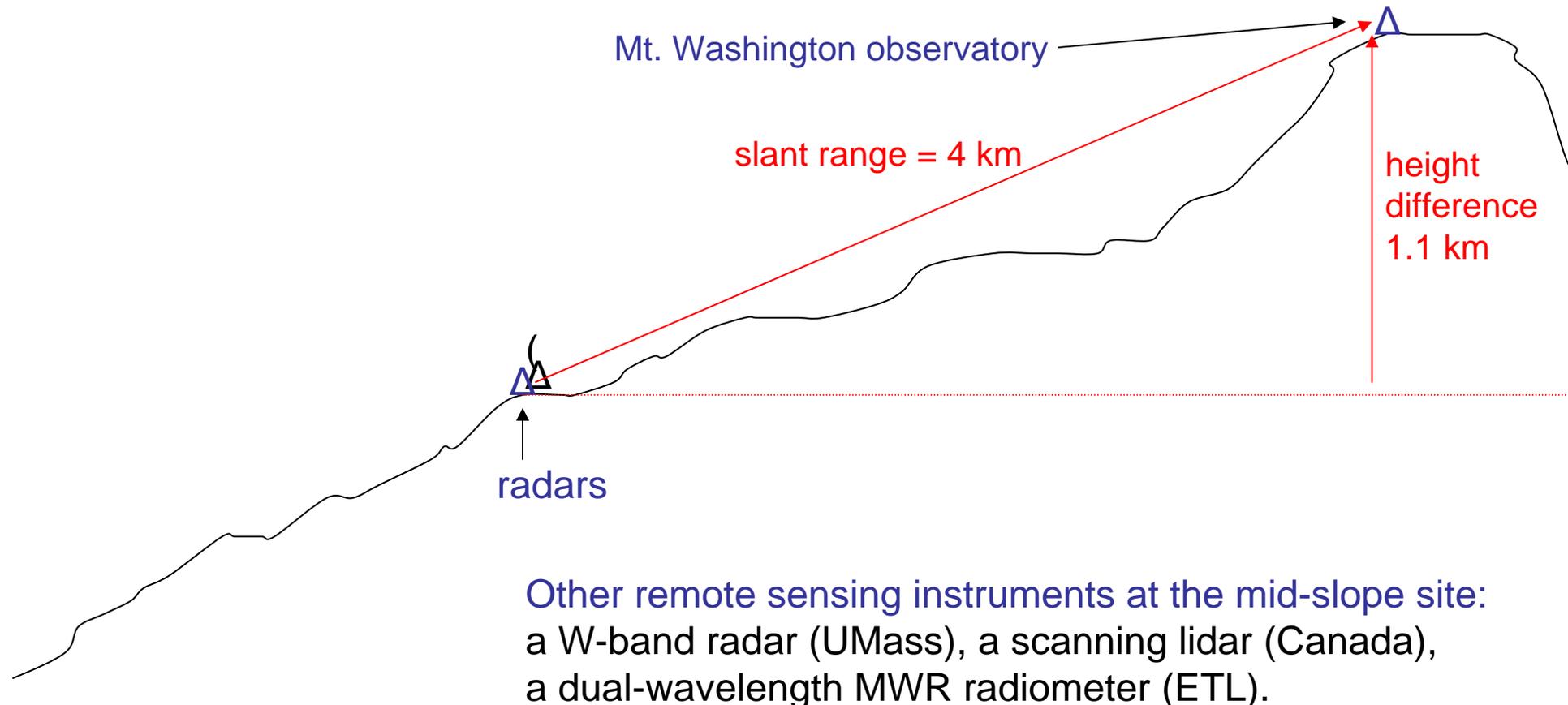
Some highlights of the MWISP field project and its relevance to StormVEx

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MWISP field experiment: April 1999, Bretton Woods, NH

MWISP = Mount Washington Icing Sensors Project

NOAA ETL deployed K_a- and X-band scanning radars at the mid slope of Mt. Washington



The main ETL's MWISP task:

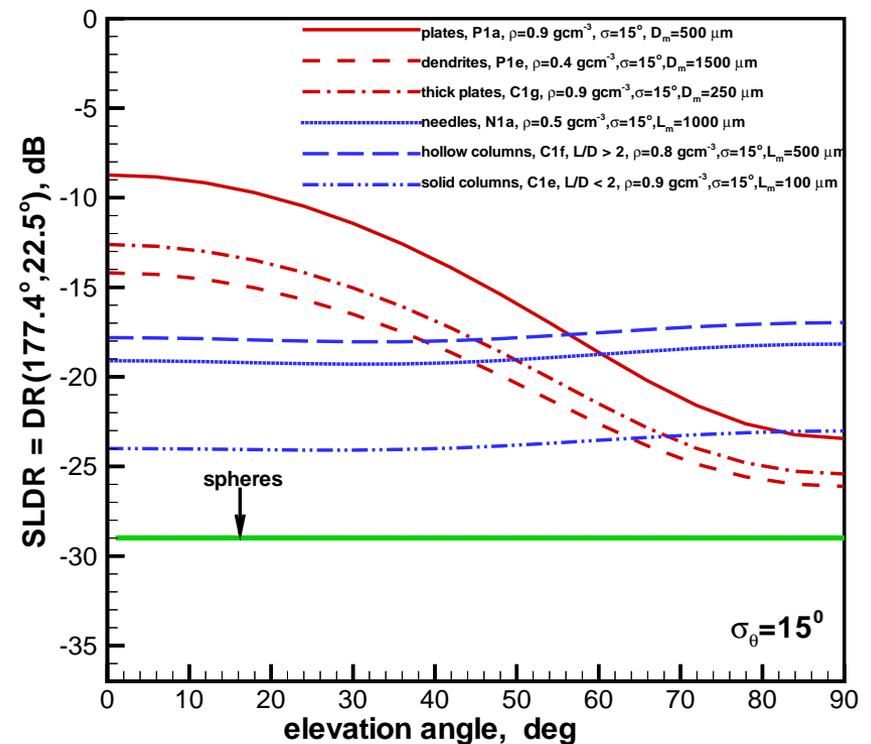
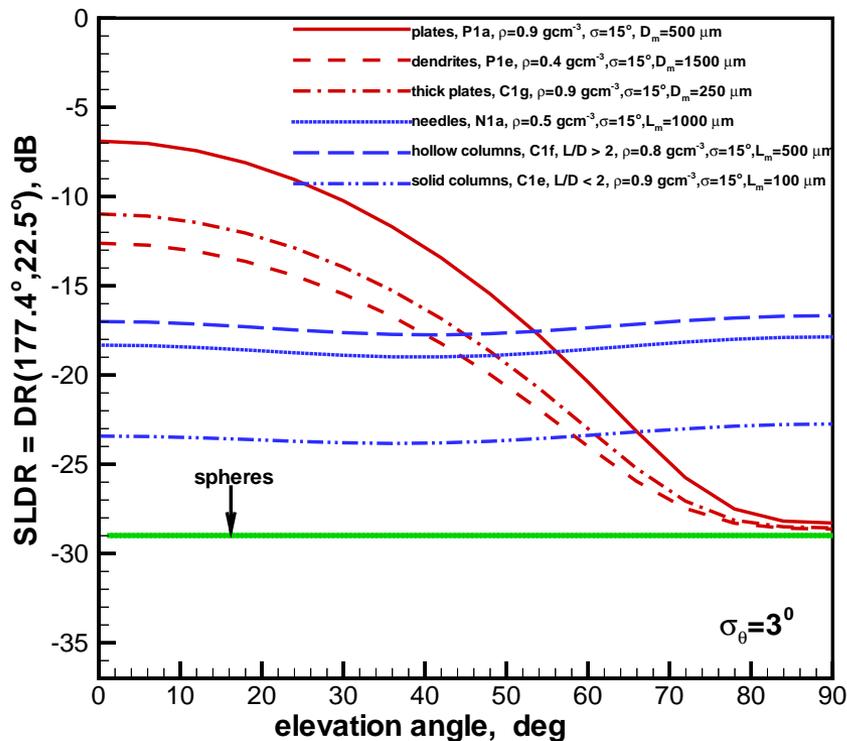
To test K_a- band polarimetric scanning radar techniques to detect supercooled large droplets (SLD).

SLD (e.g., freezing drizzle drops with sizes 30 – 400 μm) cause a significant aircraft icing hazard and usually are accompanied by few to no ice crystals

Small supercooled liquid drops that often co-exist with ice crystals are usually responsible only for minor rates of aircraft performance degradation

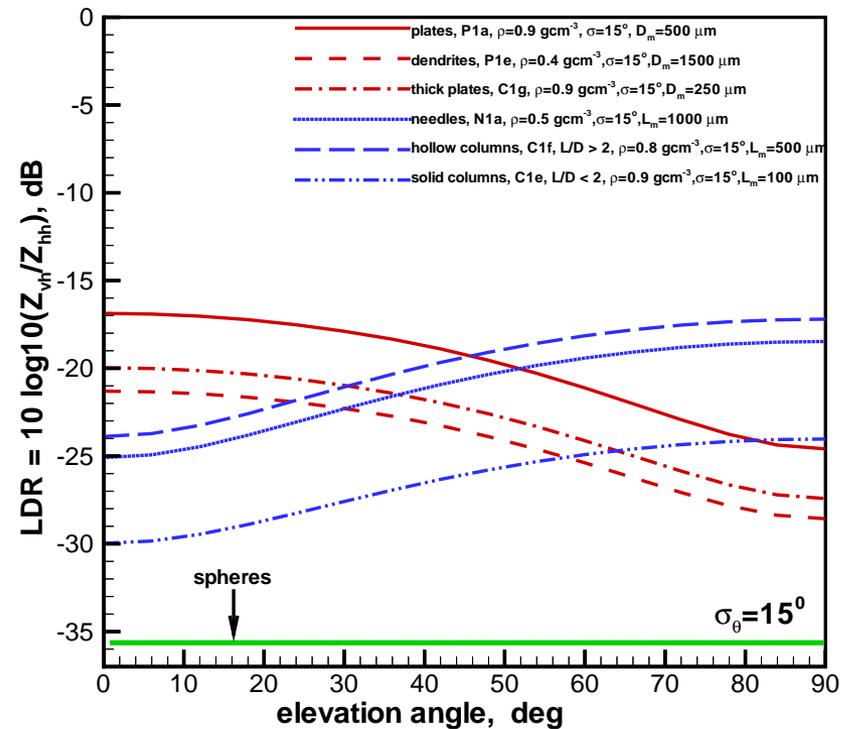
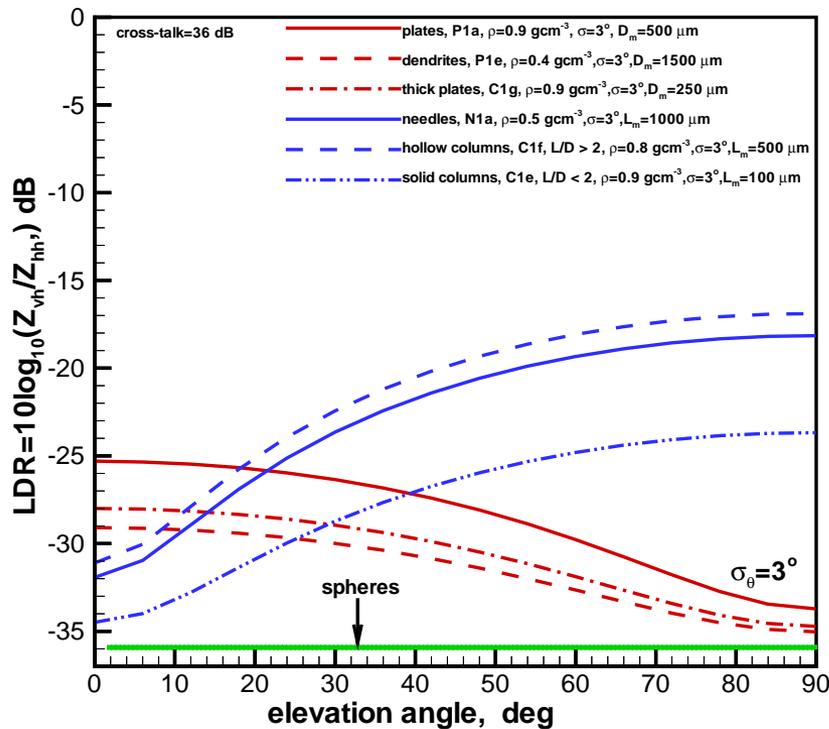
K_a -band SLDR ($SLDR=10\log_{10}(Z_{45\text{deg}135\text{deg}}/Z_{45\text{deg}45\text{deg}})$) was the polarimetric parameter used for discriminating different habits of atmospheric hydrometers

SLDR allows distinguishing among planar and columnar ice crystals and also spherical SLD when performing RHIs

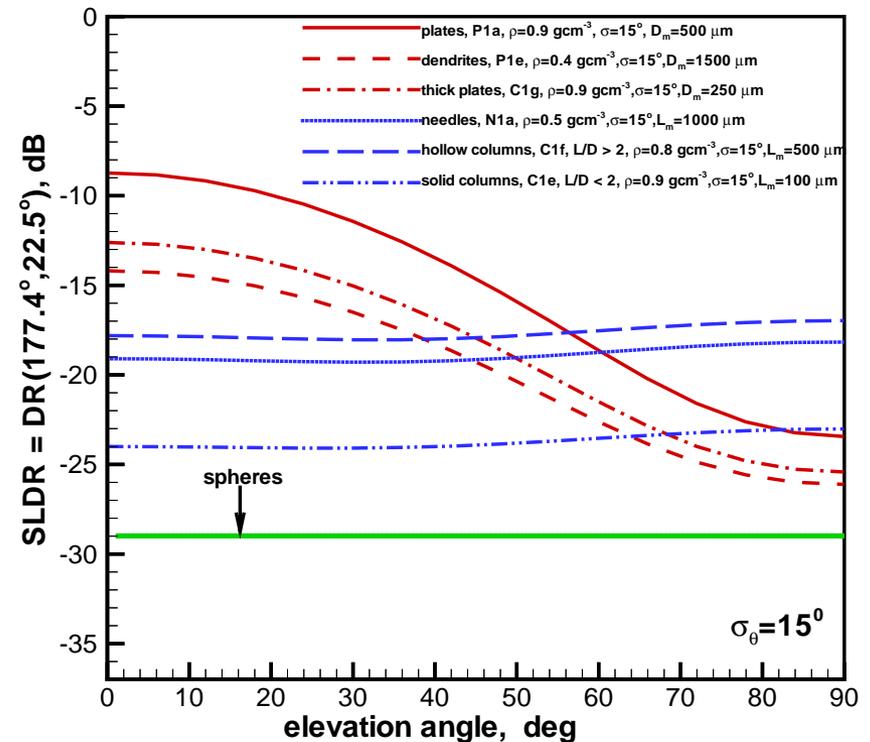
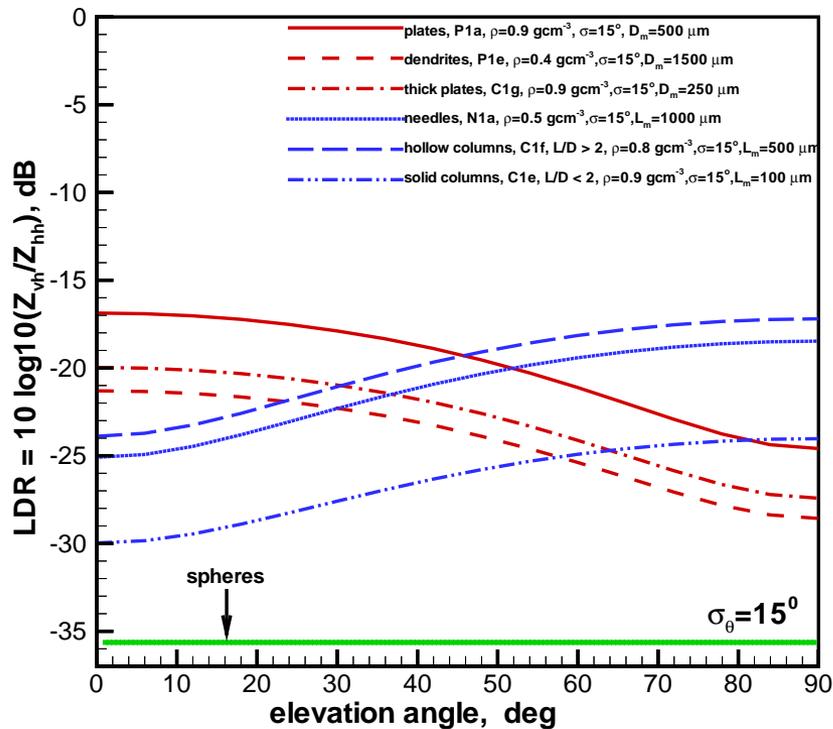


SLDR is superior to conventional H-V LDR for discriminating hydrometeor shapes

(SLDR depends on hydrometeor orientation only slightly while LDR depends on hydrometeor orientation strongly)



SLDR is superior to conventional H-V LDR for discriminating hydrometeor shapes



Depolarization by regular (pristine) ice crystals and warm drizzle (a)

SLDR – top of each frame

Blue: $SLDR < -28$ dB

Green: $-28 \text{ dB} < SLDR < -20$ dB

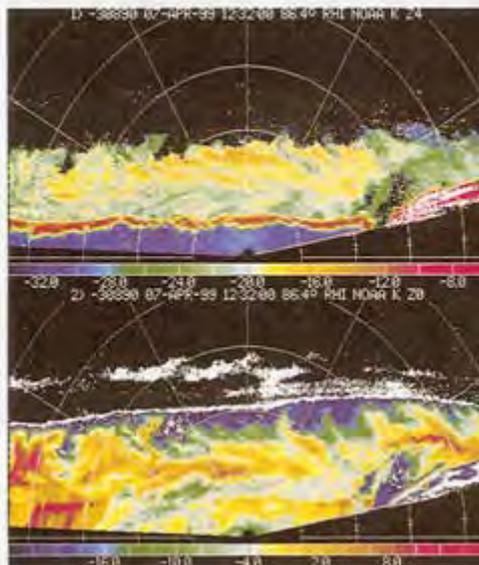
Yellow: $-20 \text{ dB} < SLDR < -12$ dB

Red : $-12 \text{ dB} < SLDR$

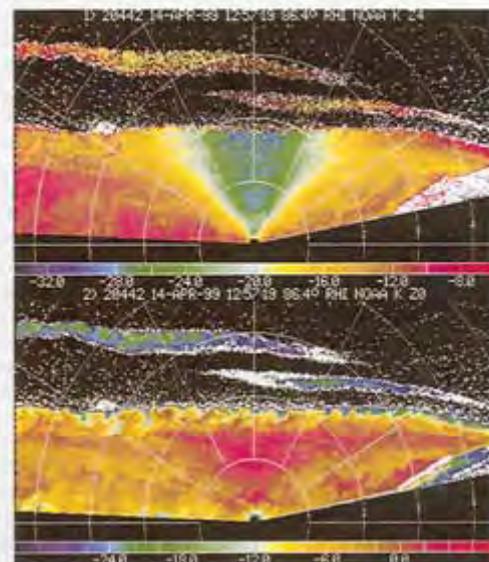
Z_e – bottom of each frame

Hydrometeor habits were verified by particle samples at the Mt. Washington

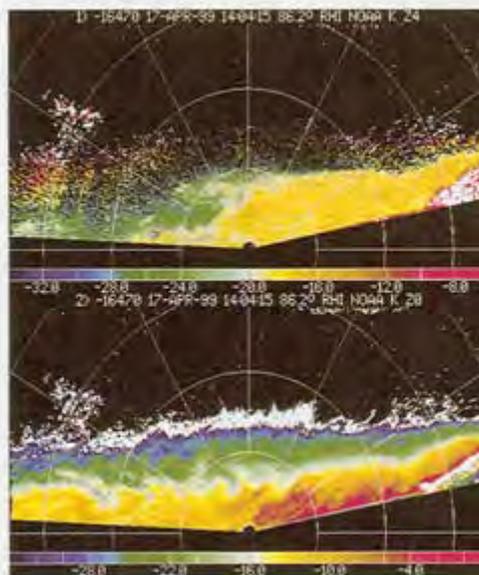
(a) ice particles aloft, melting layer, drizzle



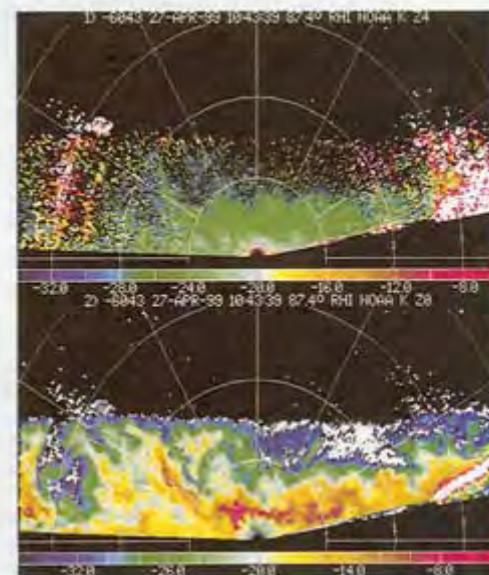
(b) dendrites



(c) long columns



(d) blocky columns



Depolarization by irregular ice crystals and supercooled drizzle (d)

SLDR – top of each frame

Blue: $SLDR < -28$ dB

Green: $-28 \text{ dB} < SLDR < -20$ dB

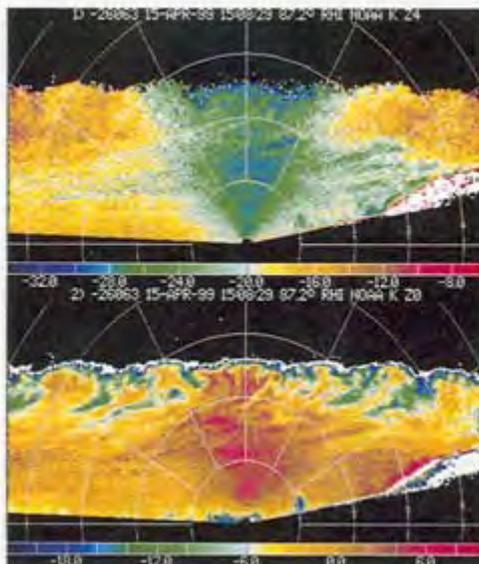
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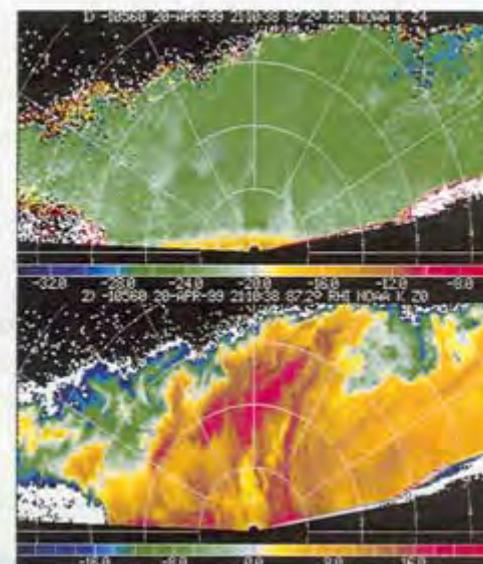
Z_e – bottom of each frame

Hydrometeor habits were verified by particle samples at the Mt. Washington

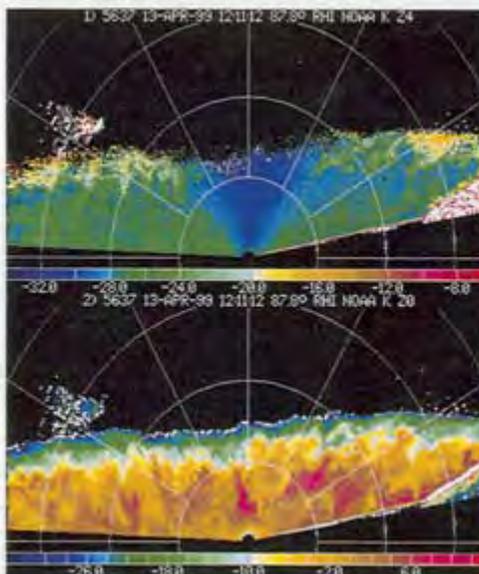
(a) aggregates of dendrites



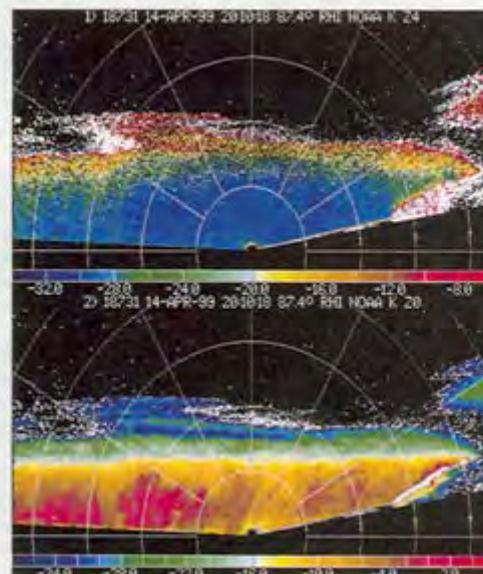
(b) conical graupel



(c) hexagonal graupel

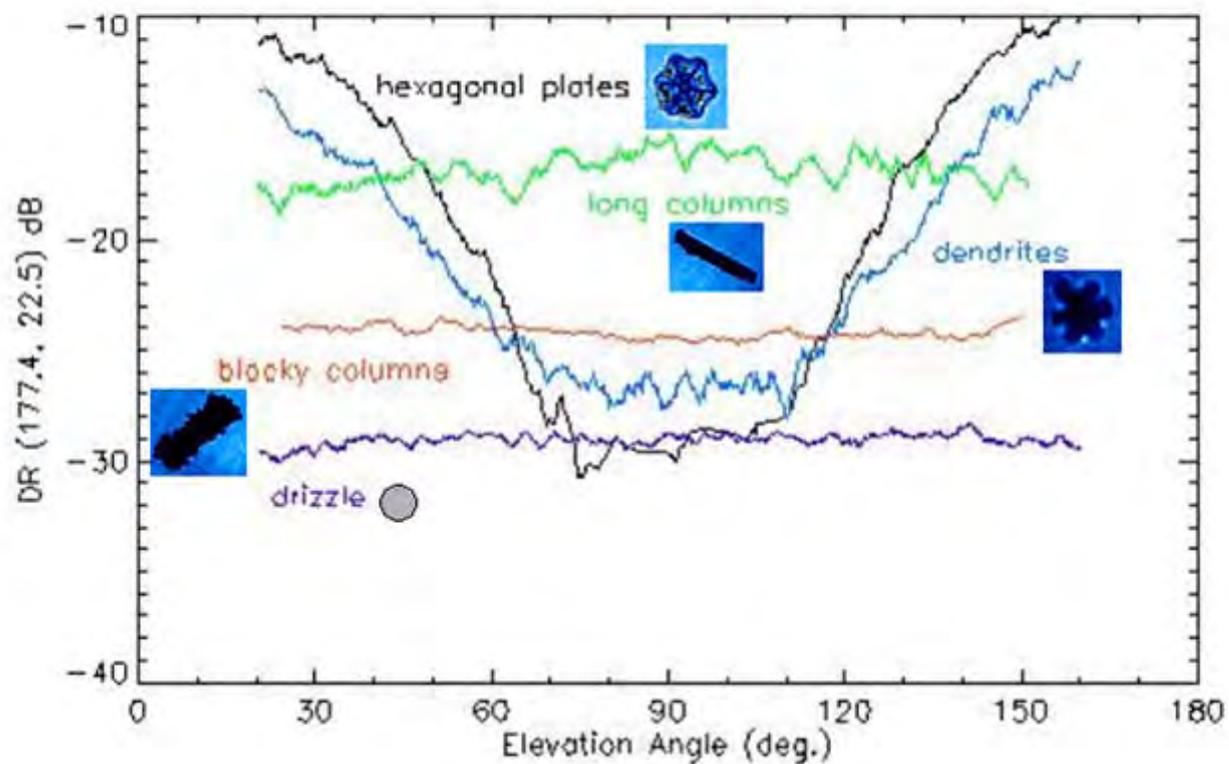


(d) supercooled large droplets

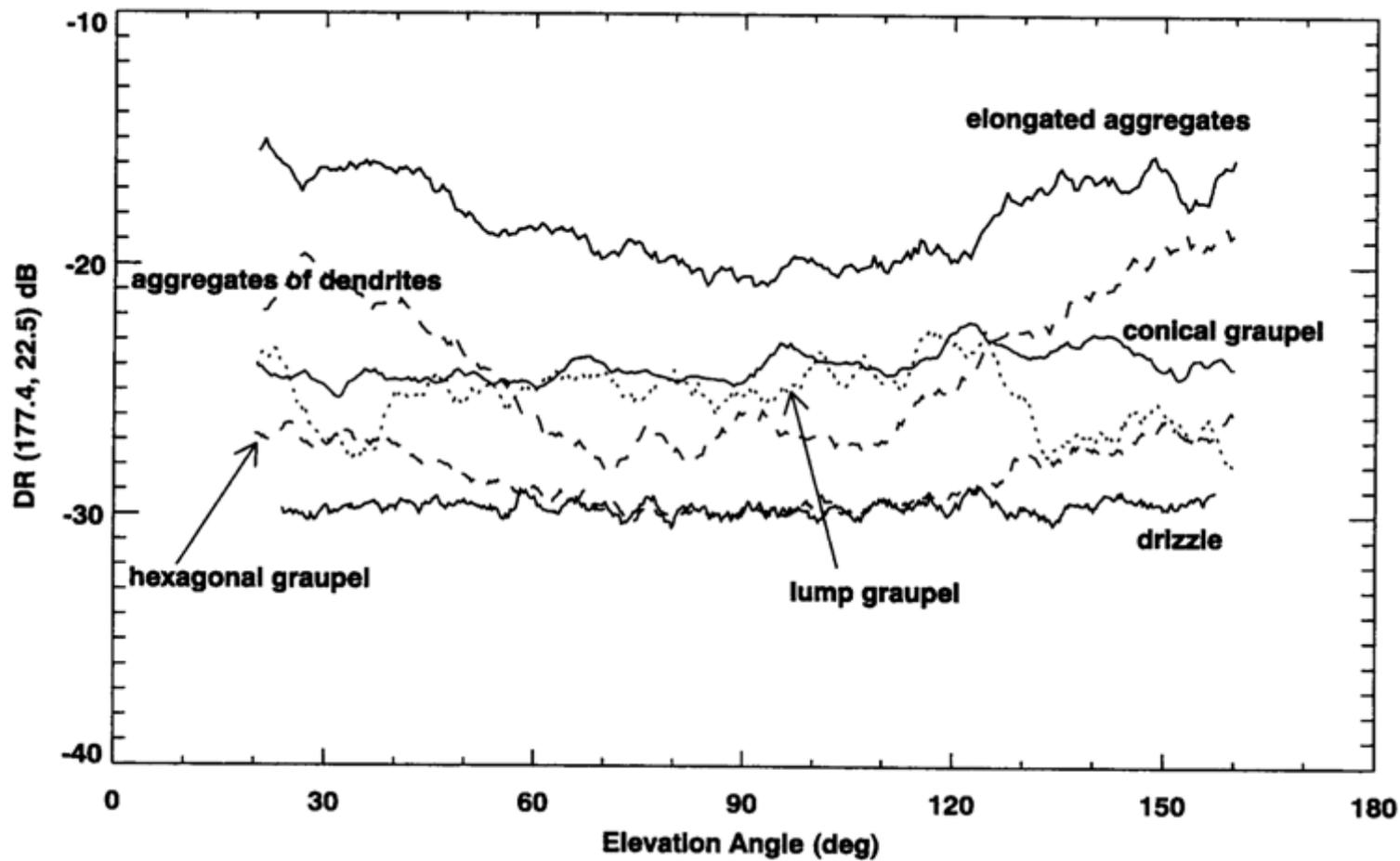


SLDR versus elevation angle for pristine crystals

45 degree slant quasi-linear polarization

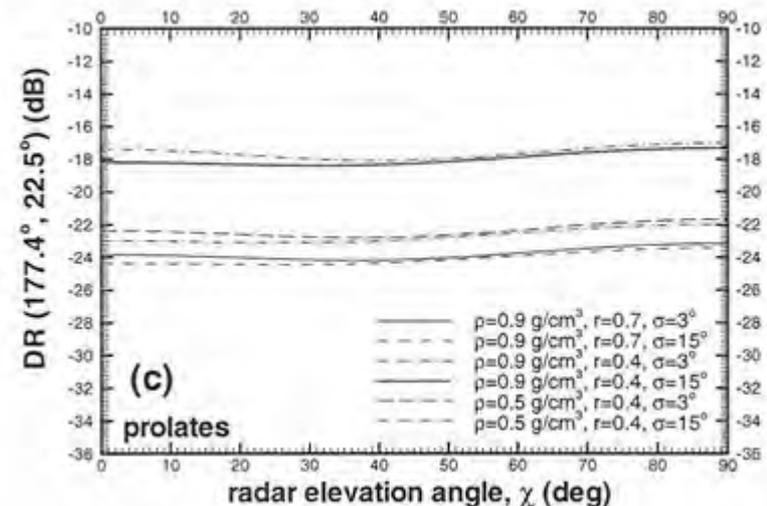
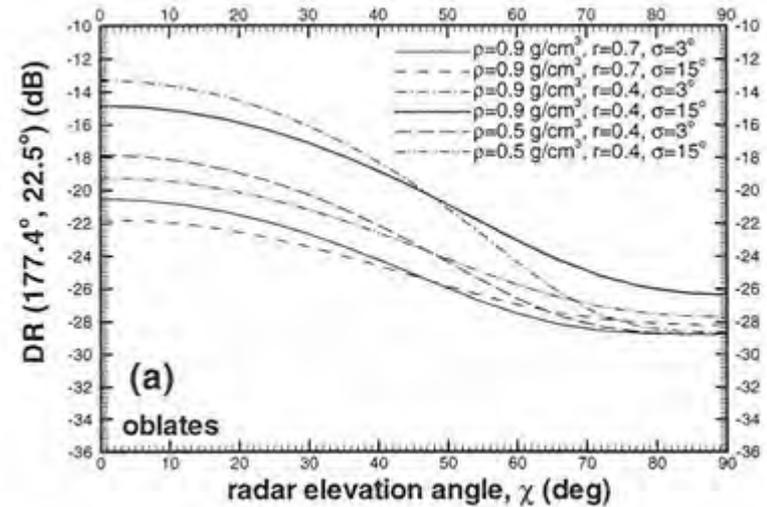


SLDR versus elevation angle for aggregate crystals



Depolarization measurements can be used not only to identify the predominant habit of ice hydrometeors but also to estimate their aspect ratio (if an assumption about the bulk density is made)

SLDR at an elevation angle of about 45° approximately does not depend on particle “wobbling” (i.e., σ_θ) and is determined for a given type (prolate or oblate) by the aspect ratio and density



A main polarimetric radar question to be addressed by StormVEx:

Could SLDR measurements at W-band (SWACR) be used for hydrometeor type identification and aspect particle ratio estimations in a way it was done at Ka-band in MWISP ?

Habit ID and particle aspect ratio retrievals can add a new dimension to ARM cloud retrievals.

Issues that would complicate SLDR signatures at W-band:

- Stronger Mie scattering effects would affect SLDR values making aspect ratio -SLDR relations less pronounced compared to K_a -band
- Stronger differential propagation effects will also affect W-band SLDR increasing the apparent (i.e., observed) SLDR values
- Non-simultaneous measurements of co-cross polar echoes by SWACR (one receiver) might negatively impact quality of SLDR measurements