Using observations of clouds and radiation at Chilbolton UK to evaluate NWP models.

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- Chilbolton 24h/7d vertical profiles of clouds
- 94GHz radar and lidar profiles 30sec/60m resolution.
- Infer cloud properties and compare with values held in operational models for Chilbolton grid box.
- 35GHz radar, 22/28/38GHz Radiometers, Raman lidar.
- 1275 clear air radar boundary layer + refractivity
- 3GHz polarisation radar for precipitation.

COMPARE OBSERVATIONS AND REPRESENTATION IN MODELS.

- Typical day
- Is cloud fraction correct? Pdf OK?
- Is ice water content correct? Errors? Pdf OK?
- Errors when classified by weather regime?
- Example of one month data and model
- Cloud overlap not really maximum random?
- Cloud inhomogeneity?
- Supercooled layer clouds are common.

Standard Chilbolton observations on the web Radar Lidar, gauge, radiometers 3 June 2003 Attenuated backscatter coefficient Radar reflectivity factor 10-3.5 10-10leight (km) Height (km) 10-10-5 - 21 101 10-6.9 00.00 03:00 06:00 21:00 00:00 00.00 03:00 06.00 09.00 15.0018.00 21.00 00.00 09:00 12:00 15:00 18:00 Time (UTC) Time (UTC) 50 Rain rate Doppler velocity 11 10 But can the average user Height (km) make sense of these 21:00 00:00 03:00 measurements? Doppler spectral width 11 Height (km) າ 250 ອີ ອີ 200 10-0. 150 10-100 50 10-1 -50 L 00:00 15:00 03:00 06.00 09.00 12:00 15:00 18:00 21.00 00.00 03.00 06.00 09.00 12.00 18.00 21.00 00.00 Radar and lidar measurements from Chilbolton Time (UTC) Time (UTC) 1000 Standard deviation of mean velocity 11 Longwave Downwelling broadband fluxes at Chilbolton 10 200 10^{-0.1} Height (km) 600 ٩ E 10-≥ 10-1.5 400 10-10-2.5 200 00:00 03:00 06:00 09:00 12:00 15:00 18:00 21:00 00:00 00:00

Time (UTC)

03:00

06:00

09:00

12:00

Time (UTC)

15:00

18:00

21:00

00:00

Target categorization

 Combining radar, lidar and model allows the type of cloud (or other target) to be identified



Cloud fraction

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Model gridboxes



Radar provides first guess of cloud fraction in each model gridbox

Lidar refines the estimate by removing drizzle beneath stratocumulus and adding thin liquid clouds (warm and supercooled) that the radar does not detect

Cloud fraction: one year of data



- Too much cloud at high levels, too little at mid-levels
 - However, frequency of occurrence is better: suggests humidity structure is good, but amount when present is not so good
 - Low-level clouds are very different in the two models

Ice water content

Met Office C-130 aircraft data



Cirrus *in situ* measurements suggest we can obtain IWC from Z to a factor of two

- Particles tend to be smaller at lower temperatures, so with additional use of temperature, error is reduced to -30%/+40%
- Less accurate between -10°C and 0°C because of strong aggregation



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    Ice water
content
from Z and T
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 Error in ice water content

- Retrieval flag
- Would be identified as ice if below freezing Clear sky above rain Ice above rain: no retrieval Ice detected only by the lidar Retrieval with correction for liquid atten. Unreliable: uncorrected attenuation Reliable retrieval No ice

Ice water content: results



- First ever long-term evaluation of ice water content
- Underestimate of mean mid-level IWC in both models
 - Seems to be due to factor-of-2 error in mean cloud fraction
 - Mean in-cloud IWC appears to be reasonably good above 4 km



IWC distributions

 The Met Office Unified Model tends to simulate very high and very low ice water contents too infrequently



Clasification by regime



- Ascent at 700 hPa (>0.1 hPa/s)
- Ascent at 400 hPa (>0.1 hPa/s)
- Stability between 900 and 1000 hPa
- Descent and low level stable UM can't make 100% cloud cover

Eddy dissipation rate *E*

 30-s standard deviation of 1-s radar velocities, plus wind speed, gives eddy dissipation rate (Bouniol et al. 2003)





PDF of & by cloud type

- Use classification product to define simple cloud types, then look at PDF of eddy dissipation rate for each
- Mean turbulence in different clouds:
 - Stratocu: 10⁻³ m²s⁻³
 - Mixed-phase: 10⁻⁵ m²s⁻³
 - Cirrus: 10⁻⁶ m²s⁻³





Radar Observations Nov 2003 Z

 σ_v

 $\sigma_{v(1sec)}$



Observations Nov 03 Lidar

Rainfall

LWP

LW and SW flux





Status

Radar ground clutter Radar corrected for liquid atten. No radar but known attenuation Good radar echo only No radar but unknown attenuation Good radar & lidar echos Radar echo but uncorrected atten. Lidar echo only Clear sky



Observations Nov 03

Cloud fraction

Total cloud cover



IWC Error

IWC

Would be identified as ice if below freezing Ice above rain: no retrieval Ice detected only by the lidar Retrieval with correction for liquid atten. Unreliable: uncorrected attenuation

Turb K E Dissipation

Nov 03 ECMWF Model



Nov 03 ECMWF model



U

V

W

Nov 03 ECMWF Model



Cloud overlap assumption in models

 Cloud fraction and mean ice water content alone not sufficient to constrain the rad



- Assumptions generate very different cloud covers
 - Most models now use "maximum-random" overlap, but there has been very little validation
 - of this assumption

Cloud overlap from radar: example



Radar can observe the actual overlap of clouds We next quantify the overlap from 3 months of data

Cloud overlap: results



- Vertically isolated clouds are randomly overlapped
- Overlap of vertically continuous clouds becomes rapidly more random with increasing thickness

• ANALYTICAL EXPRESSION FOR VERTICAL DECORRELATION Hogan and Hingworth (QJ 2000)

5. Importance of ice-cloud inhomogeneity

- Non linear relation between optical depth and emissivity
- For clouds which are inhomogeneous use of average optical depth gives wrong emissivity.



Pomroy and Illingworth (GRL 2000)

Cirrus fallstreaks and wind shear - inhomogeneities



Ice water content distributions

- In the near future, models will carry variables for the variance of water content, as well as the mean
- Derive variance of ice water content of cirrus from radar



 PDFs of IWC within a model gridbox can usually be fitted by a lognormal or gamma distribution

Analytic expression for effect of shear on pdf of iwc and vertical decorrelation as a function fo grid box size.



- Variance and decorrelation increase with gridbox size
 - Shear makes overlap of inhomogeneities more random, thereby reducing the vertical decorrelation length
 - Shear increases mixing, reducing variance of ice water content

Hogan and Illingworth (JAS 2003)

Mixed-phase clouds – SUPERCOOLED LAYER CLOUDS

- SUPERCOOLED LAYER CLOUDS ARE COMMON
- SAME WATER CONTENT BIG RADIATIVE EFFECT IF LIQUID DROPLETS – SMALL EFFECT IF ICE PARTICLES.

SUPERCOOLED CLOUD EXAMPLE

- Chilbolton 94 GHz reflectivity factor and 905 nm lidar backscatter coefficient, 21/12/98 10 94 GHz radar Height (km) 06:00 07:00 08:00 09:00 10:00 11:00 12:00 13:00 14:00 15:00 16:00 17:00 18:00 Time (UTC) 10 -50 -50 50 Lidar ceilometer -40 8 Height (km) 06:00 07:00 08:00 09:00 10:00 11:00 12:00 13:00 14:00 15:00 16:00 17:00 18:00 Time (UTC) (sterad m)-1 dB7 -40 -30 -20 -10 0 10 10-6 10-5 10-4
- Radar detects ice cloud (large D) at 5 to 6 km height
- Lidar detects highly reflecting layer at -20°C
- in-situ aircraft confirmed supercooled droplets (100 Wm⁻² impact)

Supercooled water – model comparison



- Use ground-based lidar to estimate \bullet occurrence of supercooled water layers over a 1-year period
- Around 15% of mid-level ice clouds at Chilbolton contain liquid water with optical depth > 0.7

How does this compare with the representation in the Met Office and ECMWF models?

Fraction of clouds with significant liquid 0 1 2 2 4 5 9 2 8 6 1 Independent ice & liquid 0.6 0.2 .002 004 .001 -50 -30 -100 -40 -20 Temperature (°C)

(a) Met Office Unified Model



Hogan et al. (QJ 2003)

DUAL WAVELENGTH 35 AND 94GHZ RADAR

- LIQUID WATER CONTENT THE 94GHZ RADAR IS ATTENUATED MORE THAN THE 35GHZ.
- ICE PARTICLE SIZE -

Z AT 94GHz -MIE SCATTERING Z AT 35GHZ - RAYLEIGH SCATTERING RATIO OF Z GIVES PARTICLE SIZE.

ONCE SIZE IS KNOWN CAN FIND N FROM Z, AND SO MORE ACCURATE IWC.

LIQUID WATER CONTENT PROFILE – FROM DIFFERENTIAL ATTENUATION OF Z AT 35 AND 94GHZ



ICE PARTICLE SIZE FROM 35 AND 94GHZ



25m dish: Scan on interesting days.



1275MHz clear air Refractivity

3Ghz – polarisation Precipitation radar.

1275MHz Clear air 'acrobat' radar.

- Return is from changes in refractive index turbulence on the scale of $\lambda/2$ or 11.7cm.
- Changes in the summer dominated by humidity.
- Beamwidth 0.75degs 660m at 50km range

CONVECTIVE 'DONUTS'



RHI–6 AUG-03 TOP OF THE BOUNDARY LAYER



REFRACTIVITY in the boundary layer.

- Ground clutter targets
- Round trip time changes with refractive index.
- Detect as phase change in return.
- Refractivity, N, 1ppm change in refractive index.
- $\Delta N = 1$ gives $\Delta \phi = 3$ deg/km (round trip).
- $\Delta N = 1$: $\approx 1\%$ change in RH (summer) or 1K
- Technique developed by Fred Fabrey

Refractivity change over four hours in Nov 03.



3GHz/10cm BETTER RAINRATES - ZDR AND KDP IN RAIN



Z >40dBZ In heavy rain

Big drops are oblate: ZDR>2dB In heavy rain

Extra horiz phase delay 40degs thru heavy rain

Cloud products on the web

http://www.met.reading.ac.uk/radar/cloudnet/quicklooks/

Interested in data/collaboration?

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